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# Quantifying the hazard of catastrophic rock avalanches

*Acknowledging  
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# Rock avalanches

1. **Characteristics**
2. **Size-frequency scaling**
3. **Implications for hazard assessments**



# Rock avalanches

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# Characteristics



## Rock avalanches

- Extremely rapid ( $\sim 10^1$  m s<sup>-1</sup>) flow of rock particles
- Volumes  $10^6 - 10^{10}$  m<sup>3</sup>
- Frequently derived from large rock falls and rockslides

## Debris avalanches (= volcanic rock avalanches)

- Higher content of H<sub>2</sub>O and clay minerals (hydrothermal alteration)
- Higher mobility (H/L)
- Frequently triggered by sector collapse

# Characteristics



## Rock avalanches

- Dry runout of  $10^3 - 10^4$  m requires very low (Coulomb) friction
- Rock particles are subject to dynamic fragmentation during motion
- Abrupt termination of runout process

## Debris avalanches (volcanic rock avalanches)

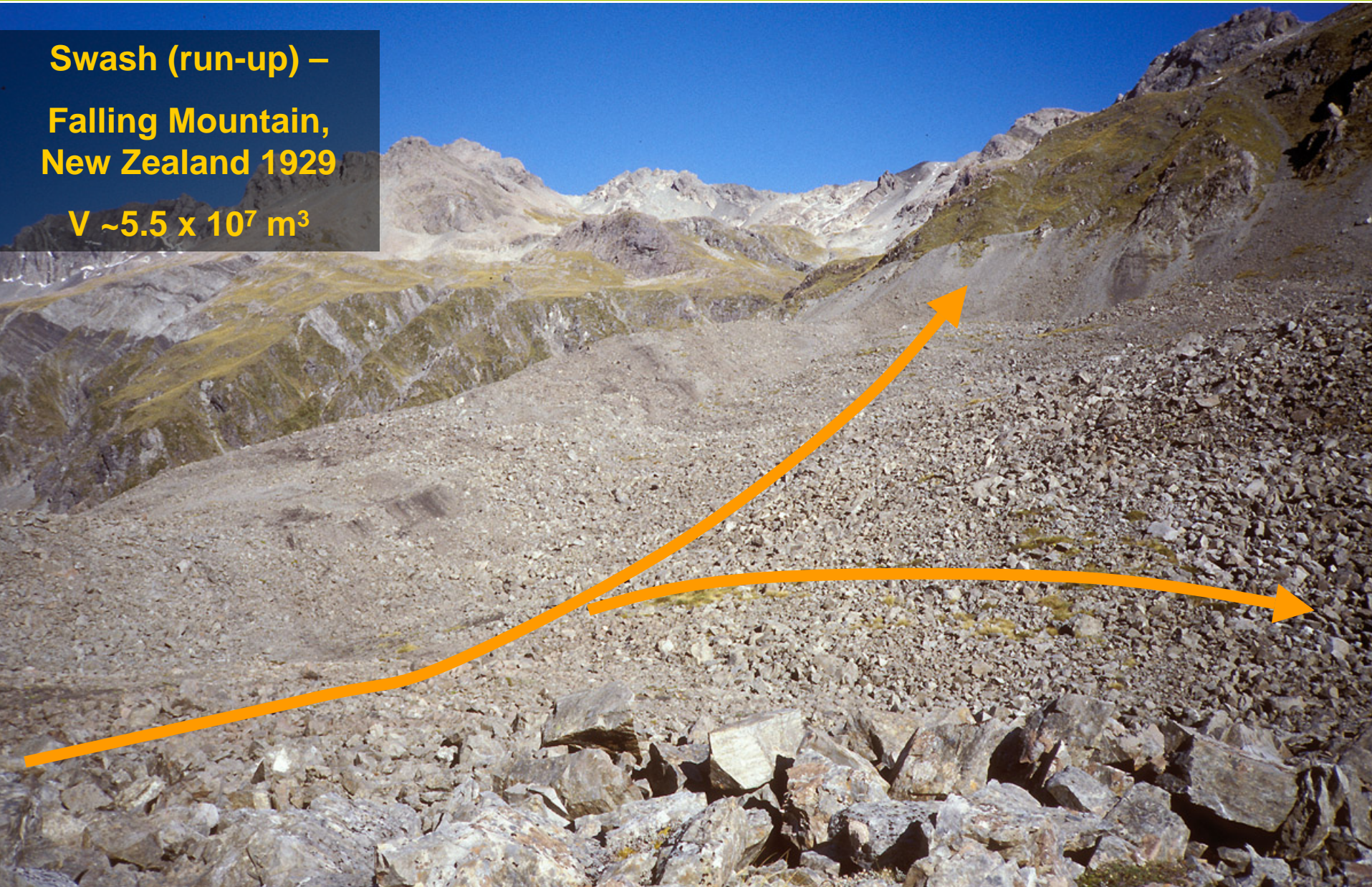
- Frequent transformation into debris flows

# Characteristics

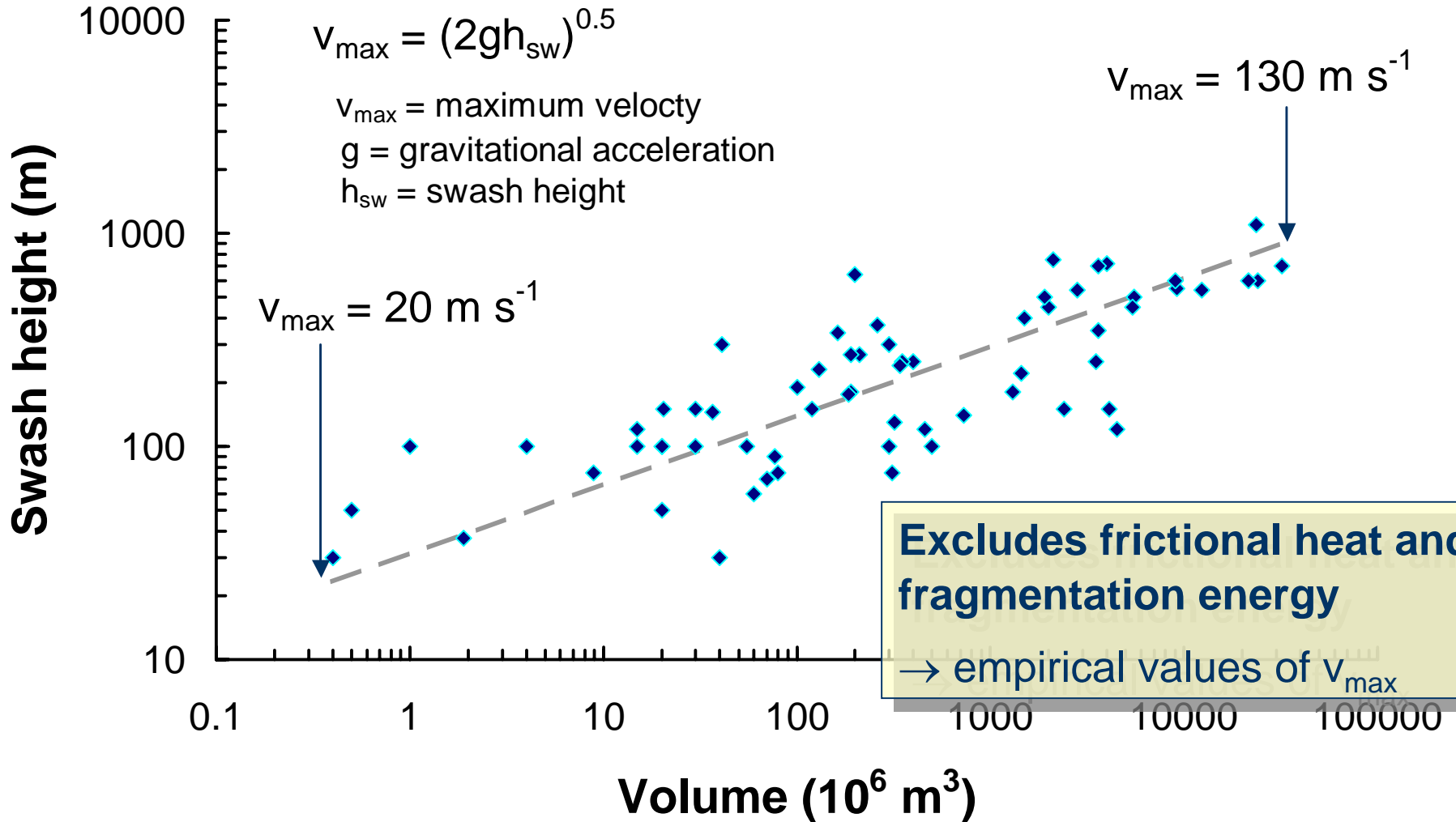
Swash (run-up) –

Falling Mountain,  
New Zealand 1929

$V \sim 5.5 \times 10^7 \text{ m}^3$



# Characteristics



# Characteristics

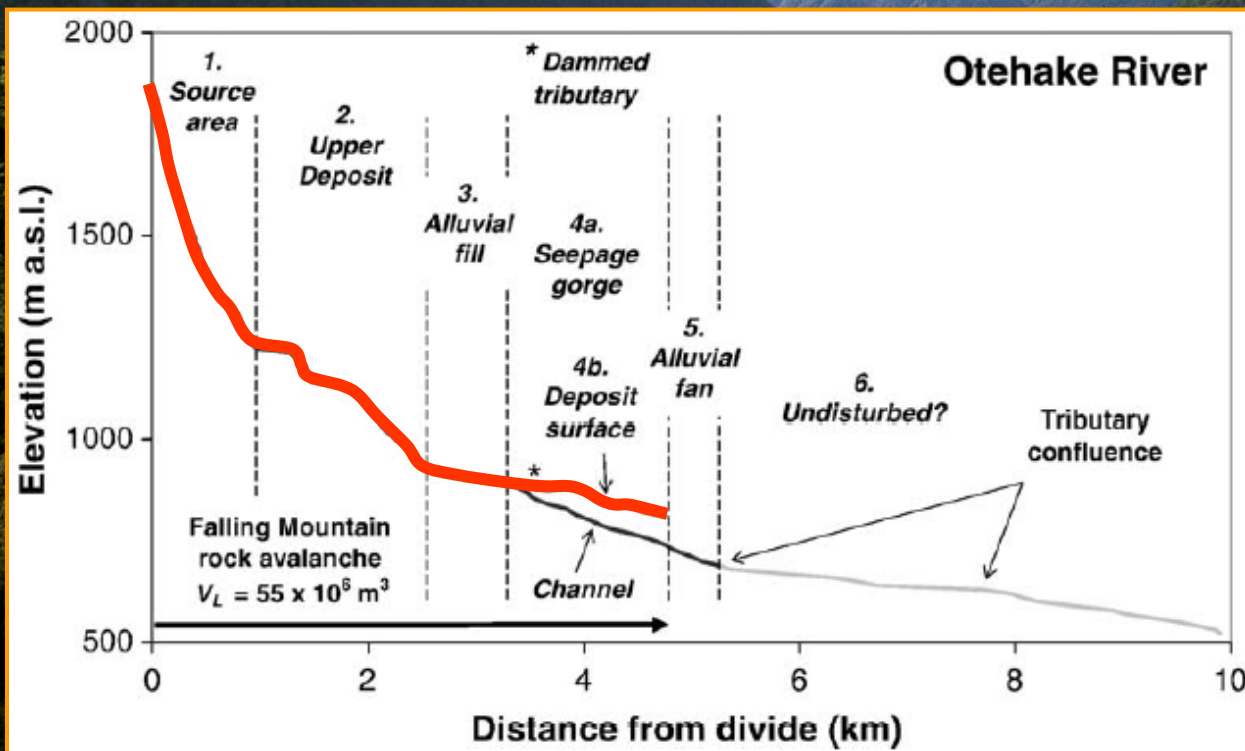
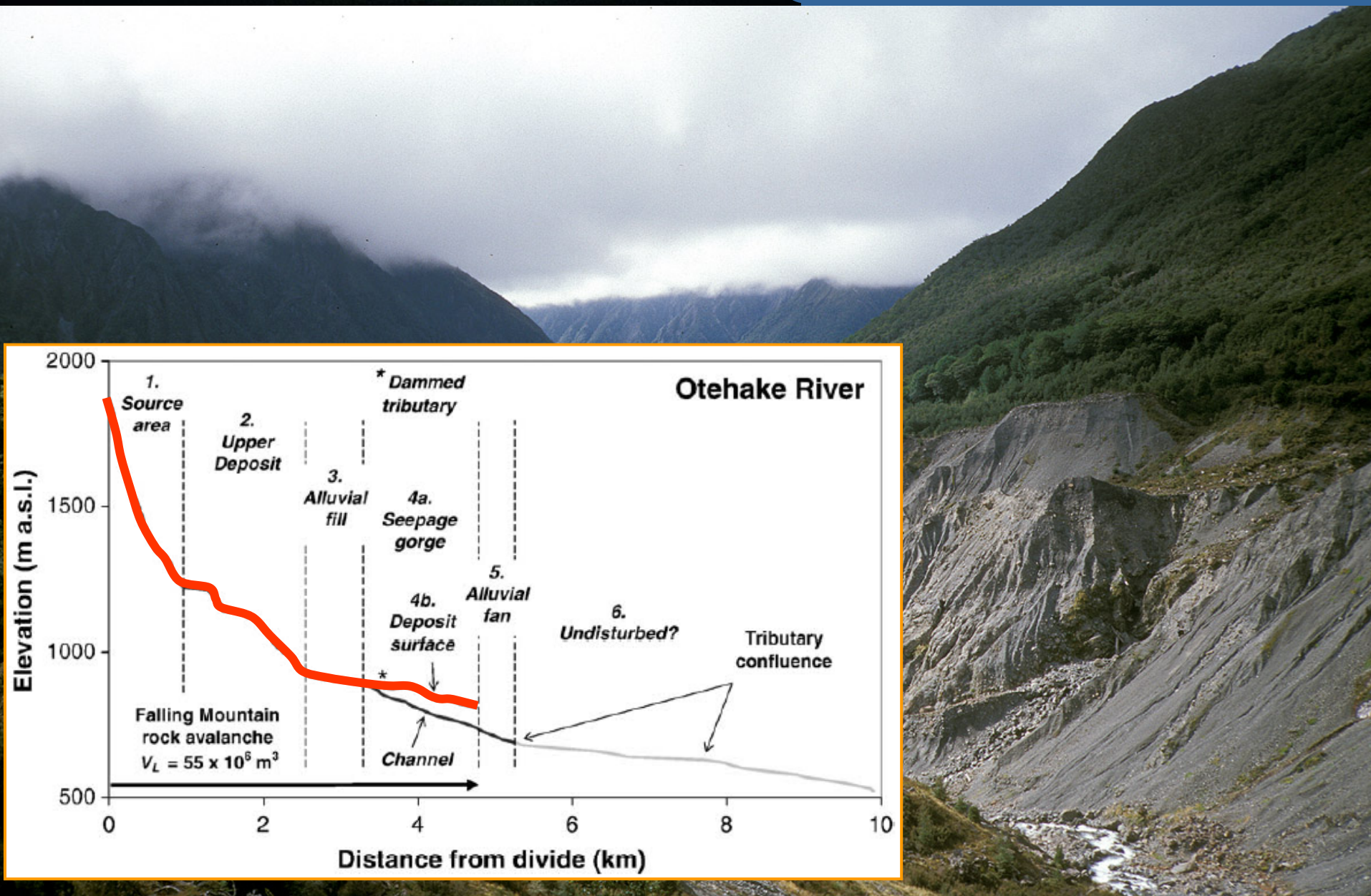




# Characteristics



# Characteristics



# Characteristics



- *Fragmentation (typically very angular clasts)*
- *Preservation of lithologic bedding (negligible vertical particle mixing)*
- *Furrows, lobes, and levees*



# Characteristics



*„Jigsaw“ puzzle texture*

## Rock avalanches

## Debris avalanches

### *Causes*

- Rock-mass discontinuities (fault zones, joints, etc.)
- Slope oversteepening (e.g. glaciation, fluvial undercutting)

- Rock-mass discontinuities (fault zones, joints, etc.)
- Hydrothermal activity
- Slope oversteepening (e.g. expansion of magma chambers)

### *Triggers*

- Earthquakes
- Rainfall and snowmelt
- Anthropogenic (mining, blasting)

- Volcanic eruption
- Sector collapse
- Earthquakes



# Rock avalanches

1. Characteristics
2. Size-frequency scaling
3. Implications for hazard assessments

- **Landslide hazard = probability of spatial landslide impact within a specified area per unit time (e.g. annual)**
- **Quantitative landslide hazard assessment thus requires a mathematical expression of the relationship between landslide size and frequency**
- **Empirical approach using records of landslide occurrence:**  
*Size-scaling relationships = statistical models that lump the complexity of spatio-temporal landslide occurrence into very few parameters*

# Quantifying hazard

## Scaling range of rock avalanches

Source: [www.gns.cri.nz](http://www.gns.cri.nz)



$10^6 \text{ m}^3$

$10^{10} \text{ m}^3$

>4 orders of magnitude



# Size-frequency scaling

- Most studies on size-frequency scaling focus on landslide area
- Largest landslide inventories contain area (volume) information on  $n \sim 10^4$  ( $n \sim 10^3$ ) data points (*though few rock avalanches...*)
- Most distributions have power-law interval  $O(2)$  to  $O(3)$

Landslide area

$$f(A) = c_A A^{-\alpha_A}$$

Landslide volume

$$f(V) = c_V V^{-\alpha_V}$$

- Little is known about the scaling properties of large catastrophic landslides ( $A > 1 \text{ km}^2$ )
- Distributions fitted to smaller landslides imply inverse power law

# Size-frequency scaling

## Most popular distributions

- Power law
- Double Pareto
- Weibull
- **Inverse gamma**

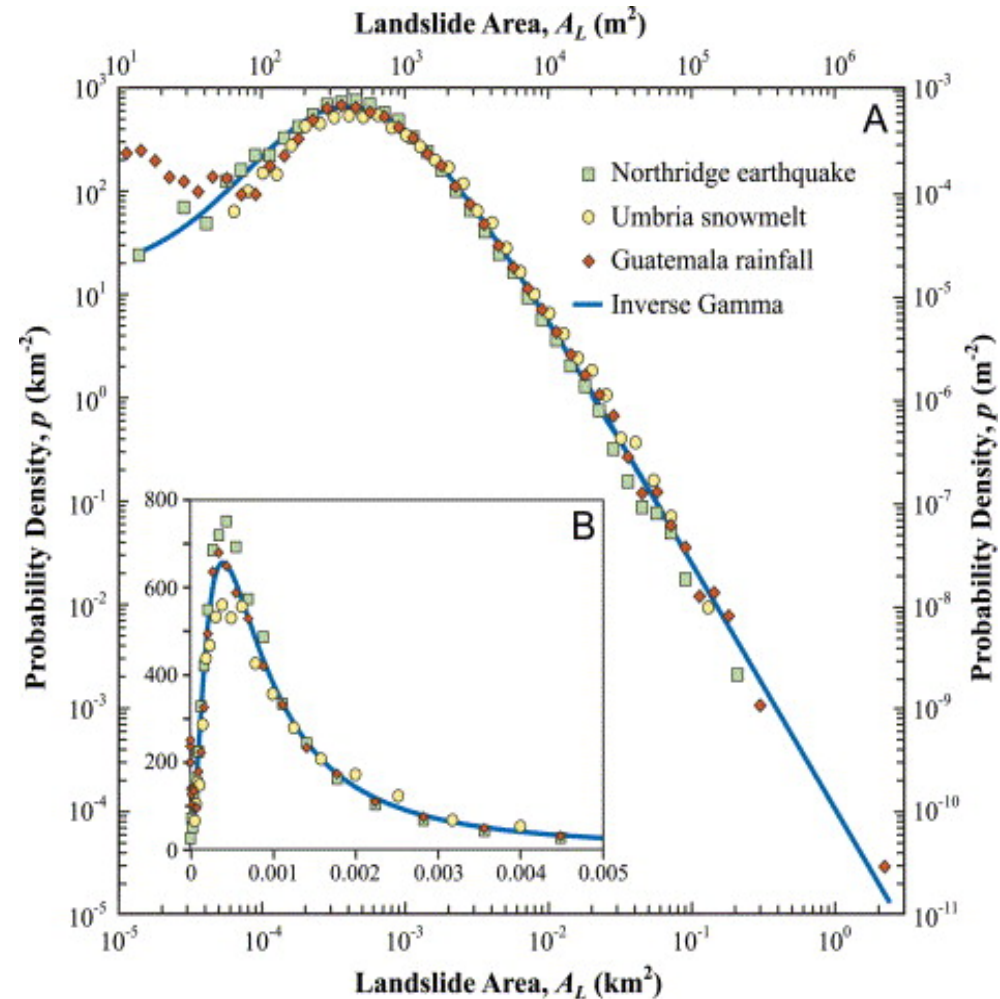
$$p(A_L; \rho, a, s) = \frac{1}{a\Gamma(\rho)} \left[ -\frac{a}{A_L - s} \right]^{\rho+1} \exp \left[ -\frac{a}{A_L - s} \right]$$

$A_L$  = Landslide area

$\rho = 1.40$

$a = 1.28 \times 10^{-3} \text{ km}^2$

$s = -1.32 \times 10^{-4} \text{ km}^2$



# Size-frequency scaling

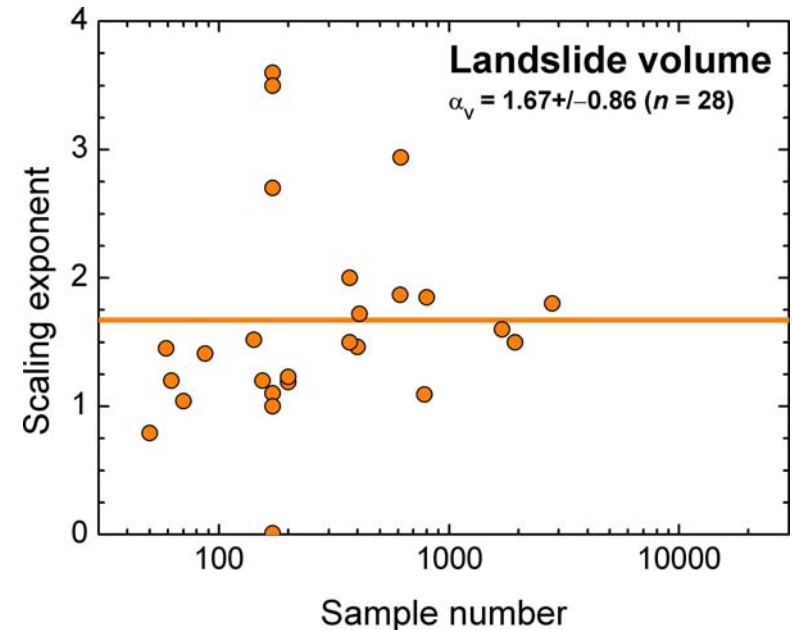
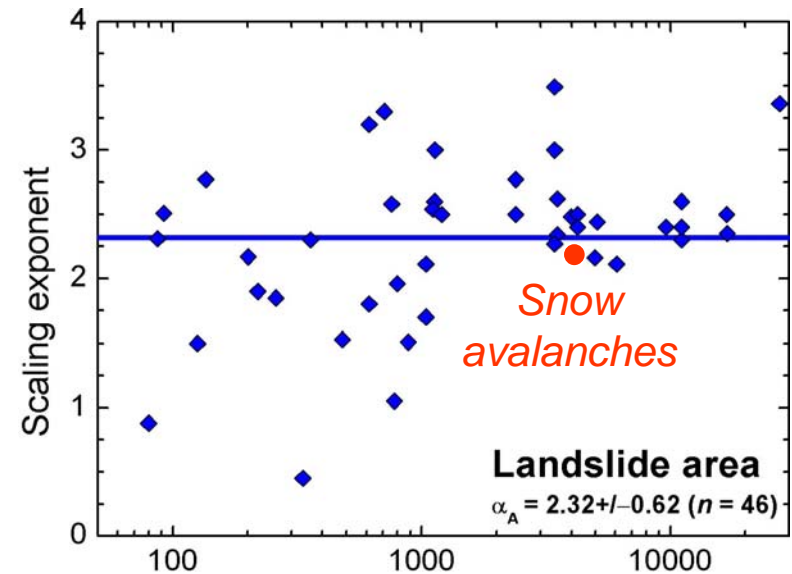
## Estimated range of scaling exponents

$$f(A) = c_A A^{-\alpha_A}$$

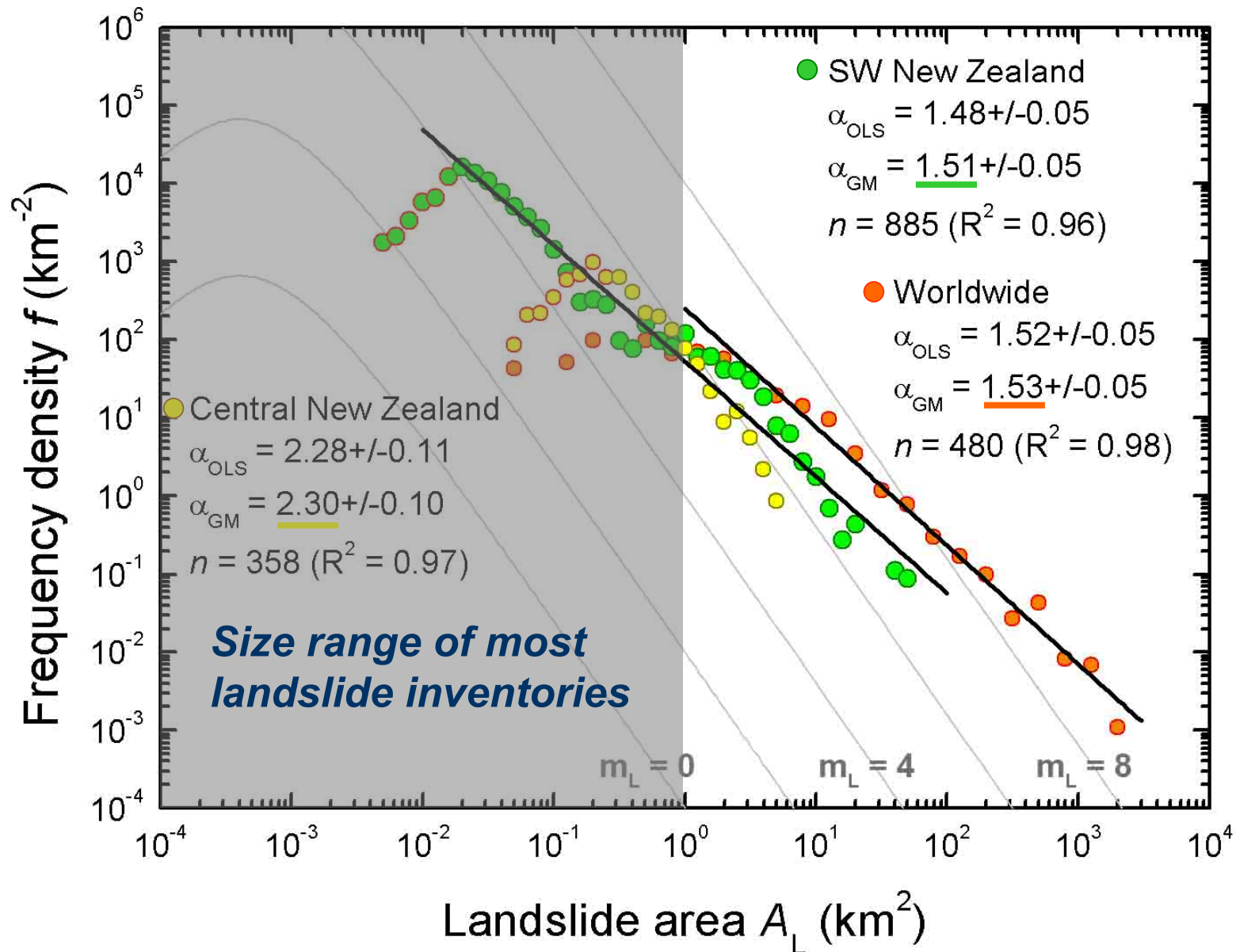
where  $2.2 < \alpha_A < 2.4$   
( $\mu \pm 1$  s.e.)

$$f(V) = c_V V^{-\alpha_V}$$

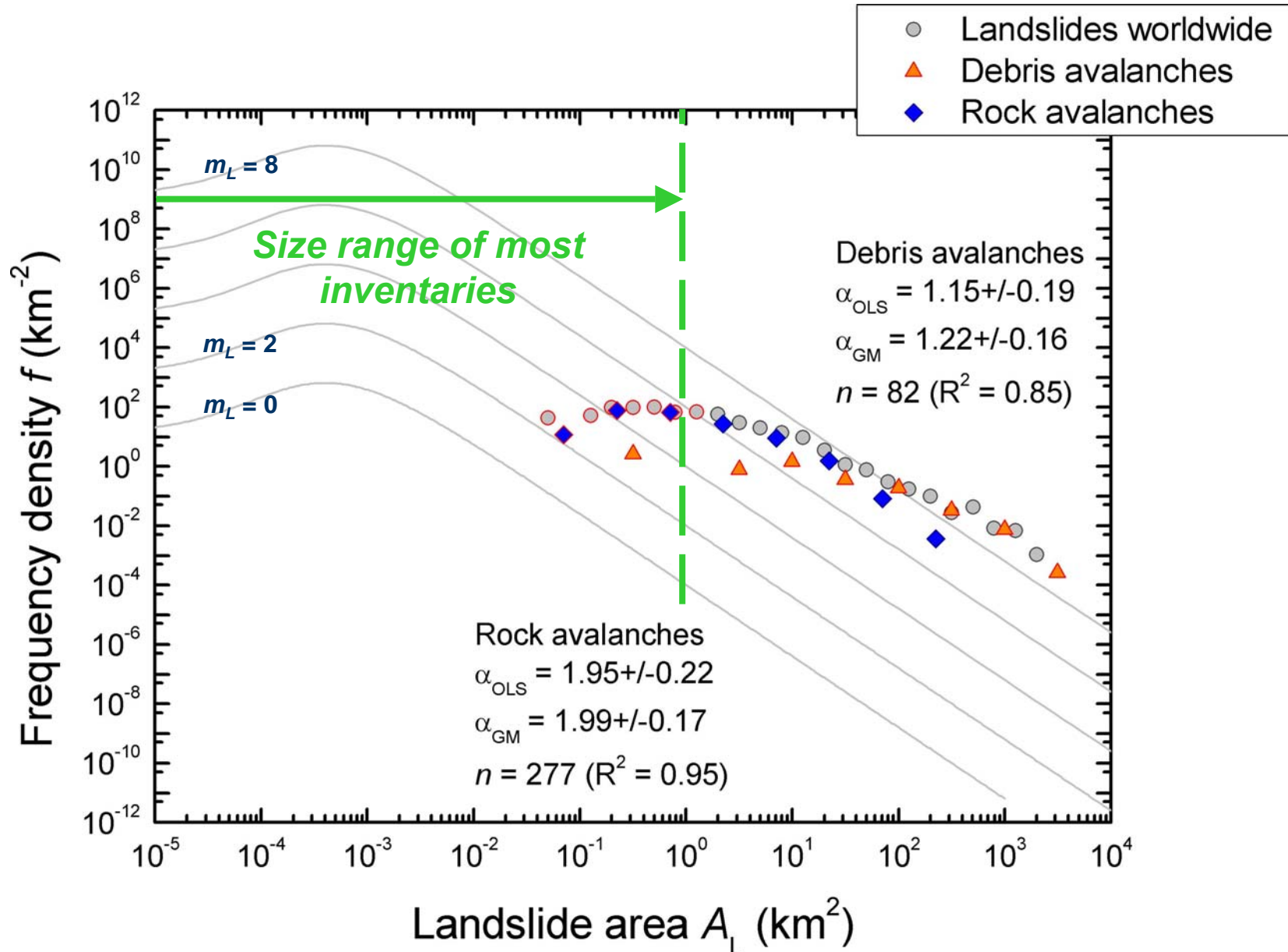
where  $1.6 < \alpha_V < 1.9$   
( $\mu \pm 1$  s.e.)



# Size-frequency scaling



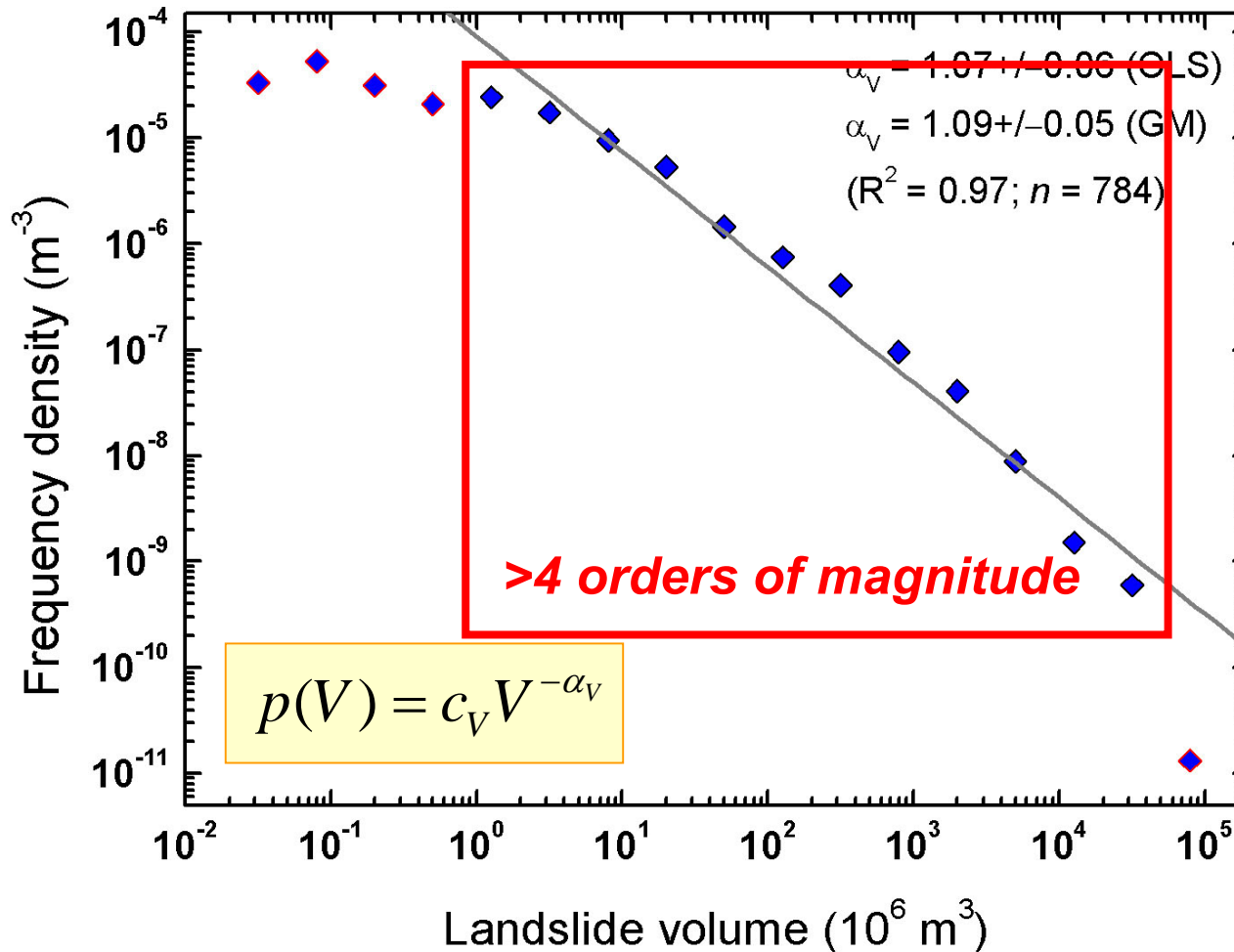
# Size-frequency scaling



# Size-frequency scaling

## Scaling properties of volume (up to $10^{10}$ m<sup>3</sup>):

*Catastrophic rock-slope failures (mainly rock and debris avalanches)*



## Estimates of scaling exponents depend on

- *Sample size*
- *Bin size*
- *Regression method (OLS, GM, Max Likelihood)*
- *Lithology*
- *Landslide type*

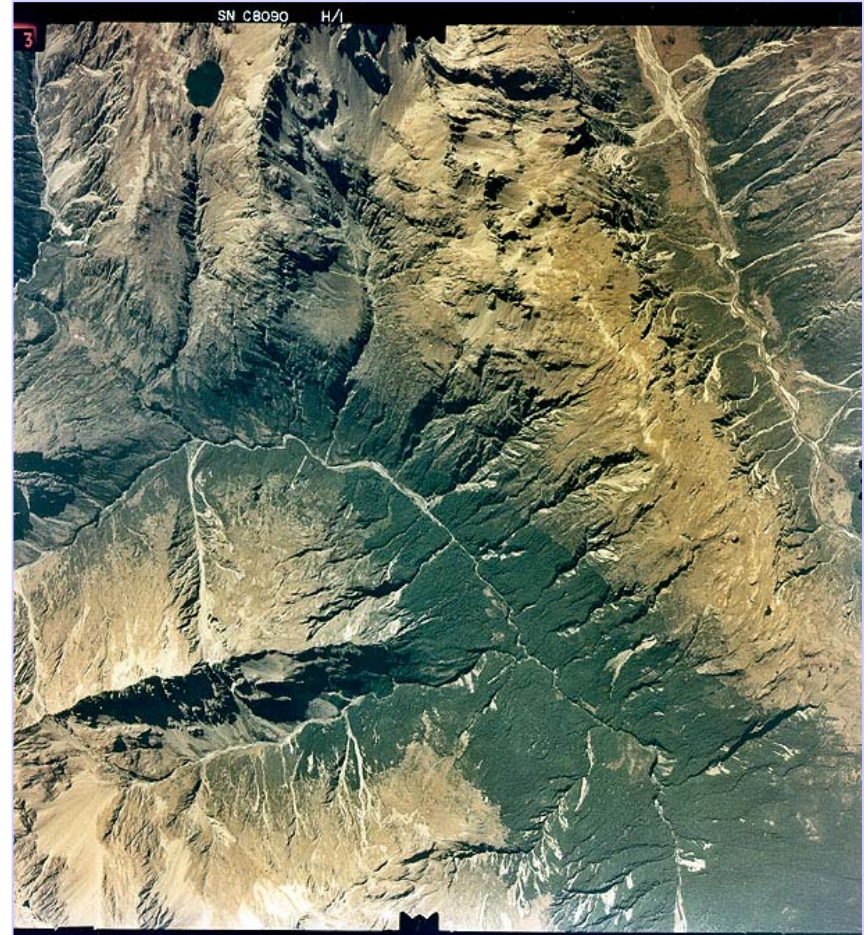
## Additional scatter comes from

- *Map projection (alpine topography)*
- *Subjective mapping*
- *Use of scar, deposit, or total affected areas*

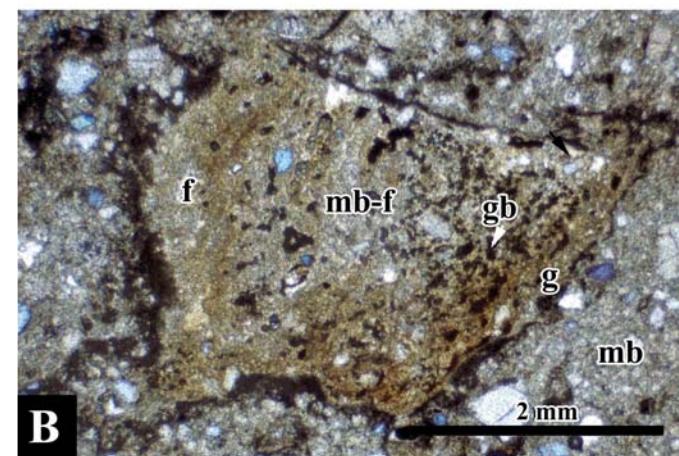
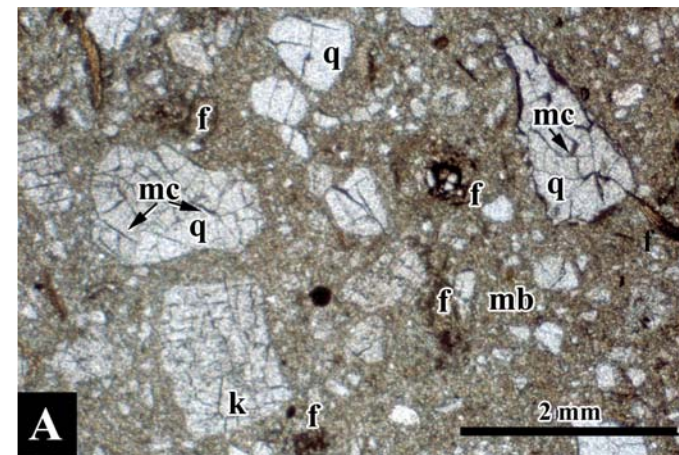
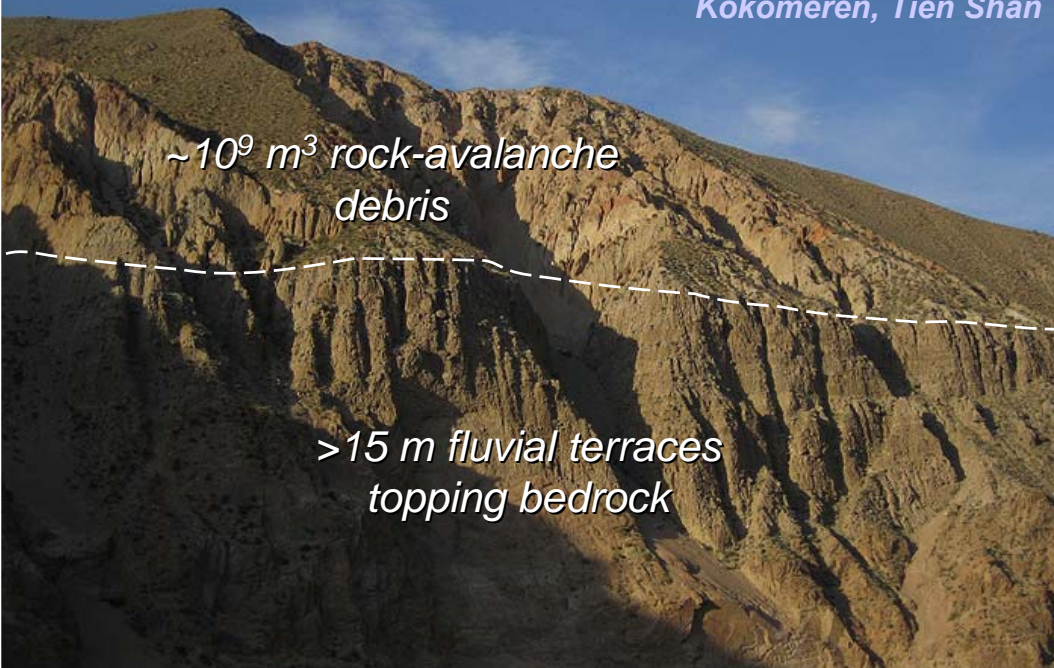
# Potential pitfalls

## ***Detection of rock avalanches using remote sensing data is limited by***

- ***Undersampling***: dense vegetation cover may mask deposits from rock avalanches
- ***Erosional censoring***: older rock-avalanche deposits may have been largely eroded
- ***Depositional censoring***: older rock-avalanche debris may have been buried by subsequent failures at the same site (reactivation)







*Frictionite (= hyalomylonite) in thin section*

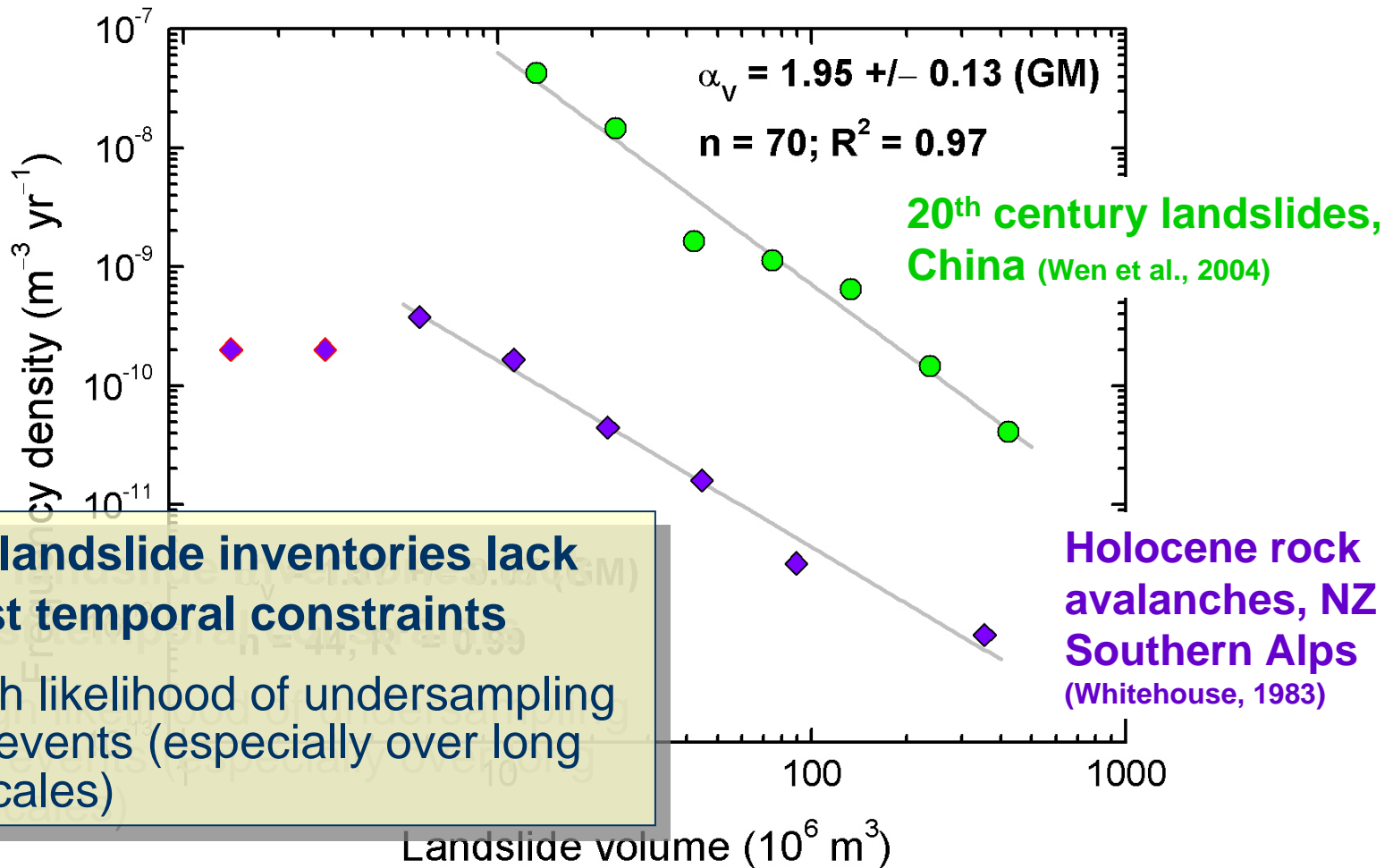
*~600 °C for partial melting of biotite*



# Rock avalanches

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# Temporal constraints

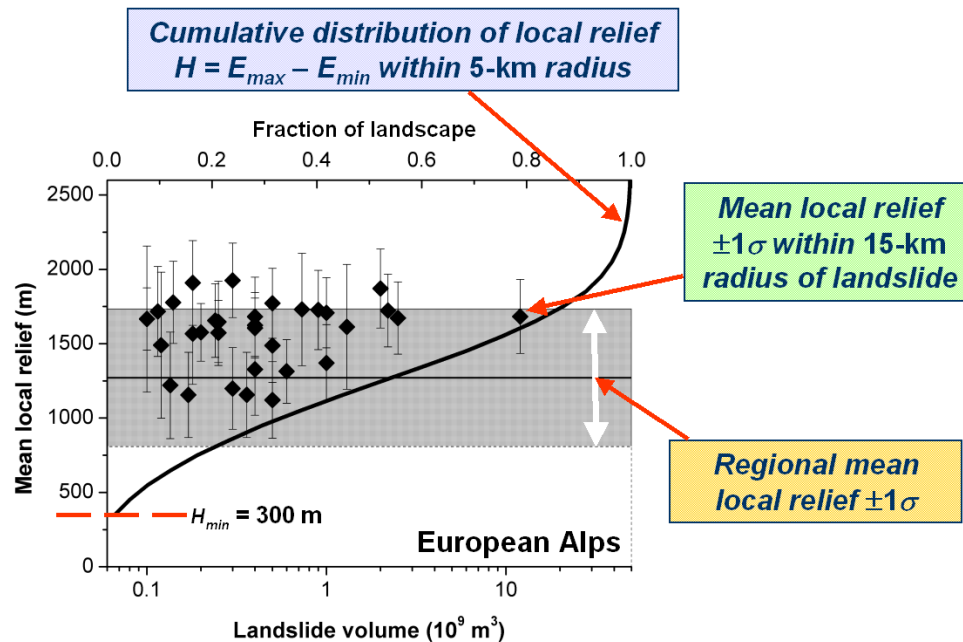


**Most landslide inventories lack robust temporal constraints**

→ high likelihood of undersampling large events (especially over long timescales)

# Spatial constraints

- Scaling-derived hazard estimates lump all spatial information and thus do not predict where a rock avalanche of a given size will occur
- Topography can be used as a first-order predictor of rock-avalanche occurrence, although failure often obliterates the geometric initial conditions



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Earth and Planetary Science Letters 261 (2007) 578–589

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## Giant landslides, topography, and erosion

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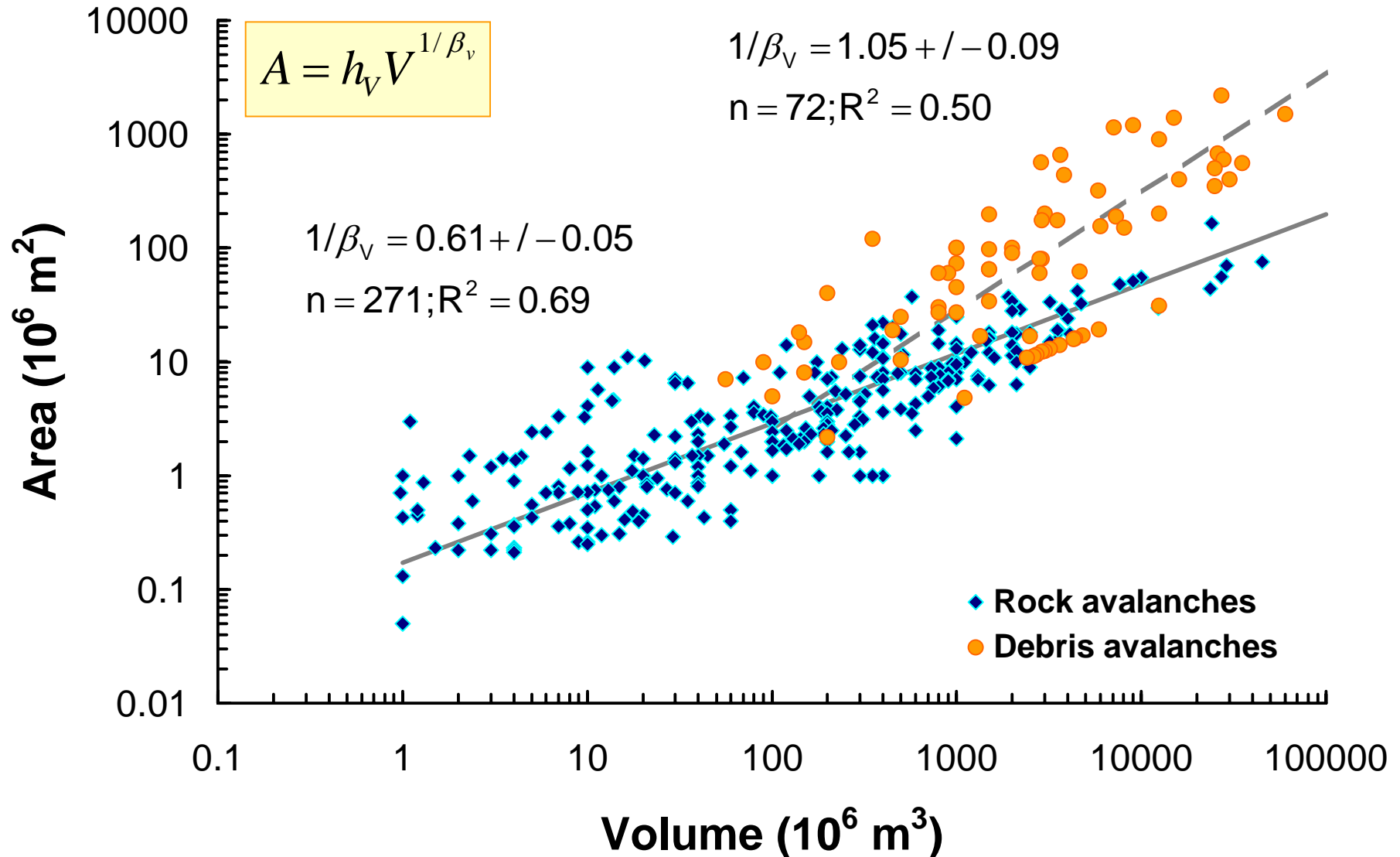
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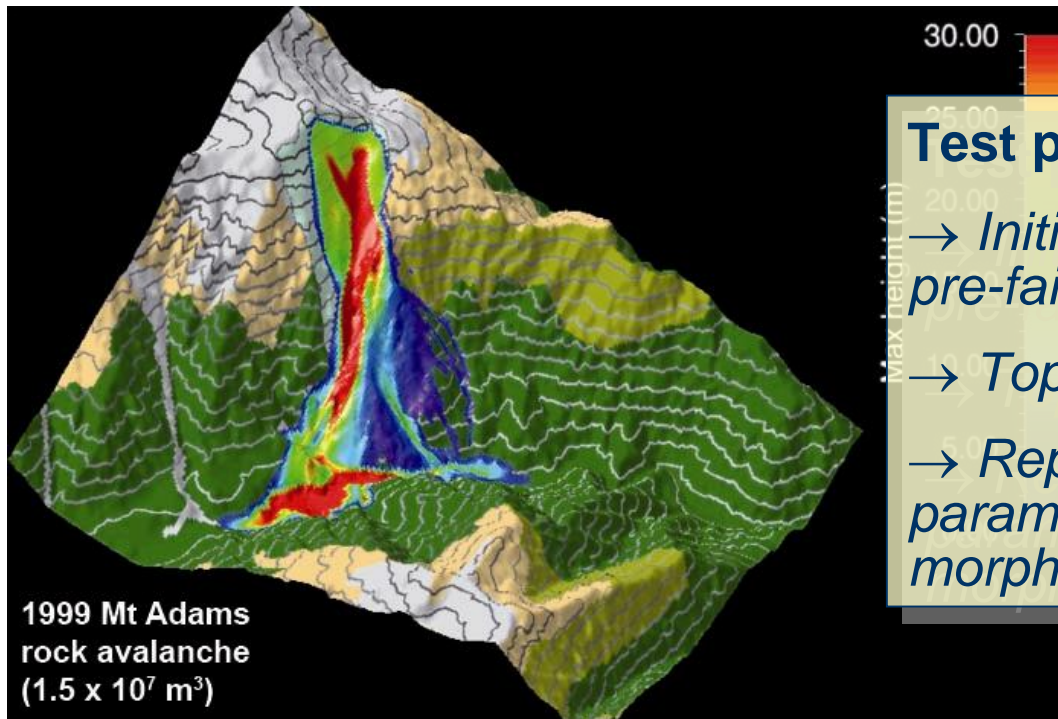
Available online 24 July 2007

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# Topographic constraints



- **Modified runout model used for snow avalanches and debris flows yields promising results for rock-ice and rock avalanches** (*i.e. the IRASMOS spectrum of extremely rapid mass movements*)
- **This model additionally considers random kinetic energy and particle collisions**



**Test phase: Model sensitivity to**

- *Initial conditions (volume and pre-failure geometry)*
- *Topography in runout zone*
- *Representation of new parameters in deposit surface morphology*

# Off-site hazards



**Young River, New Zealand, 2007**

Formation of natural dam upstream of rock-avalanche dam



*Helicopter*

## **Natural dams introduce**

→ *Off-site hazards (inundation, catastrophic dam break and outburst flows)*

→ *Potential for a cascade of mainly hydrological hazards*

The use of scaling properties to quantify the hazard of catastrophic rock avalanches promises several prospects and pitfalls

- ➔ *Area-frequency scaling of rock avalanches is well constrained by the power-law tail of an inverse gamma distribution. This trend is valid over >3 orders of magnitude, although exponents ( $1.2 < \alpha_A < 2.0$ ) are lower than those for smaller landslides ( $2.2 < \alpha_A < 2.4$ ).*
- ➔ *This scaling relationship offers a regional-scale measure of quantifying to first order the hazard from rock avalanches, given that temporal constraints are tight enough.*



➔ *Existing numerical models for simulating snow avalanches and debris flows show promising results for simulating rock avalanches. The use of these models appears to be limited more by the choice of appropriate initial conditions than parameter values.*

- ➔ **Recommendations:**
- *Invest more resources to better constrain the initial conditions of large rock avalanches.*
  - *Consider using a harmonized upscaling approach based on detailed case studies.*